



River-discharge effects on United States Atlantic and Gulf coast sea-level changes

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Edited by Anny Cazenave, Center National d'Etudes Spatiales (CNES), Toulouse, France, and approved June 5, 2018 (received for review March 29, 2018)

Identifying physical processes responsible for historical coastal sea-level changes is important for anticipating future impacts. Recent studies sought to understand the drivers of interannual to multidecadal sea-level changes on the United States Atlantic and Gulf coasts. Ocean dynamics, terrestrial water storage, vertical land motion, and melting of land ice were highlighted as important mechanisms of sea-level change along this densely populated coast on these time scales. While known to exert an important control on coastal ocean circulation, variable river discharge has been absent from recent discussions of drivers of sea-level change. We update calculations from the 1970s, comparing annual river-discharge and coastal sea-level data along the Gulf of Maine, Mid-Atlantic Bight, South Atlantic Bight, and Gulf of Mexico during 1910–2017. We show that river-discharge and sea-level changes are significantly correlated ($p < 0.01$), such that sea level rises between 0.01 and 0.08 cm for a 1 km³ annual river-discharge increase, depending on region. We formulate a theory that describes the relation between river-discharge and halosteric sea-level changes (i.e., changes in sea level related to salinity) as a function of river discharge, Earth's rotation, and density stratification. This theory correctly predicts the order of observed increment sea-level change per unit river-discharge anomaly, suggesting a causal relation. Our results have implications for remote sensing, climate modeling, interpreting Common Era proxy sea-level reconstructions, and projecting coastal flood risk.

coastal sea level | coastal river plumes | coastal flood risk | climate modeling | physical oceanography

Predicting regional sea-level changes and their coastal impacts is a grand challenge in climate research (1). To improve projections of future sea-level changes, it is important to understand the physical process responsible for past coastal sea-level changes in historical observations and proxy reconstructions. The tide-gauge record on the United States Atlantic and Gulf coasts has received considerable attention in this regard. Partly motivated by projections of rapid future sea-level rise along parts of this coastline (2, 3) and observations suggesting that this region is a hotspot of ongoing regional sea-level acceleration (4, 5), recent studies explored myriad processes influencing Atlantic and Gulf coast sea level over interannual to multidecadal time scales, including changes in the overturning circulation, Gulf Stream, Sverdrup transport, remotely generated planetary waves, along-shore winds, barometric pressure, groundwater extraction, dam retention, vertical land motion, and melting of land-based ice (6–14).

Conspicuously absent from these discussions is consideration of variable river discharge into the coastal ocean and its impact on sea level at interannual and longer time scales. More generally, global ocean circulation models often omit boundary forcing by year-to-year changes in river discharge (15). Despite being an important driver of circulation in the coastal ocean (16–24), river discharge is often overlooked as a driver of sea-level change.

Coastal currents and river plumes can be strongly trapped to the coast, making such features difficult to observe with conventional satellite altimetry and difficult to resolve in global models. The extent to which coastal sea level as observed by tide gauges reflects the influence of river discharge thus remains to be rigorously determined.

In 1971, Meade and Emery (25) compared annual river-discharge and sea-level data over the Gulf of Maine, Mid- and South Atlantic Bights, and Gulf of Mexico during 1931–1969. Using linear regression and correlation analysis, they suggested that river discharge explained 20% to 31% of detrended annual sea-level variance, such that sea level rose between 0.01 and 0.05 cm for a 1 km³ increase in annual river discharge, depending on region. Meade and Emery (25) reasoned that the sea-level rise per unit river-discharge increase was inversely related to total discharge but unrelated to the discharge per unit length of coastline. While these conclusions suggest that recent studies of United States Atlantic and Gulf coast sea level (and global ocean models more generally) are overlooking a potentially important driver of coastal sea-level change, many questions remain open. Meade and Emery (25) alluded to rapid dynamic adjustment and horizontally uniform spreading of fresh water over the shelf in their interpretation of the observed relations between river discharge and sea level along the coast. However, they provided no formal physical framework within which to interpret their mainly statistical findings, so it remains to be determined whether their results reflect correlation or

Significance

River discharge exerts an important influence on coastal ocean circulation but has been overlooked as a driver of historical coastal sea-level change and future coastal flood risk. We explore the relation between observed river discharge and sea level on the United States Atlantic and Gulf coasts over interannual and longer periods. We formulate a theory that predicts the observed correspondence between river discharge and sea level, demonstrating a causal relation between the two variables. Our results highlight a significant but overlooked driver of coastal sea level, indicating the need for (1) improved resolution in remote sensing and modeling of the coastal zone and (2) inclusion of realistic river runoff variability in climate models.

Author contributions: C.G.P., A.C.K., and S.E.E. designed research; C.G.P. and K.B. performed research; C.G.P. and K.B. analyzed data; and C.G.P., A.C.K., R.M.P., C.M.L., S.E.E., and S.J.L. wrote the paper.

The authors declare no conflict of interest.

This article is a PNAS Direct Submission.

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This article contains supporting information online at www.pnas.org/lookup/suppl/doi:10.1073/pnas.1805428115/-DCSupplemental.

Application of Theory to Observations

The form of Eq. 6 is consistent with basic results here and in Meade and Emery (25). That is, there is no functional dependence on coastline length, and $\partial\eta_0/\partial Q_F$ varies with the inverse (square root) of Q_F , such that larger regression coefficients are anticipated for regions with smaller river discharge and vice versa. To determine more rigorously whether this theory is consistent with observed relations between river discharge and sea level (Figs. 1 and 2), we compare estimates of $\partial\eta_0/\partial Q_F$ determined empirically from ordinary linear regression of the data to those predicted by Eq. 6. We use observed Q_F values (Fig. 1), f values computed based on the average latitudes of the tide gauges for a region, a standard value of $g = 10 \text{ m s}^{-2}$, and representative values of $S_0 = 35$ practical salinity units (PSUs) and $\beta = 8 \times 10^{-4} \text{ PSU}^{-1}$.

We compute theoretical $\partial\eta_0/\partial Q_F$ values of $0.031 \text{ cm}\cdot\text{y}\cdot\text{km}^{-3}$ for the Gulf of Maine, $0.021 \text{ cm}\cdot\text{y}\cdot\text{km}^{-3}$ for the Mid-Atlantic Bight, $0.023 \text{ cm}\cdot\text{y}\cdot\text{km}^{-3}$ for the South Atlantic Bight, and $0.007 \text{ cm}\cdot\text{y}\cdot\text{km}^{-3}$ for the Gulf of Mexico (Fig. 2). For the Gulf of Mexico, theoretical and observational $\partial\eta_0/\partial Q_F$ values are very similar, whereas for other regions, theoretical values are smaller than the corresponding best estimates from observations (Fig. 2). However, for all regions, theoretical and observational values are of the same order of magnitude, and theoretical $\partial\eta_0/\partial Q_F$ values are within the estimated observational uncertainties (Fig. 2). This suggests that observed relations between river discharge and coastal sea level are causal and interpretable in terms of basic physical principles (i.e., salt and volume conservation, geostrophic and hydrostatic balances).

However, the fact that the theoretical regression coefficients consistently underestimate observed best estimates hints that additional considerations are needed to fully understand the data. One possibility is that, while we used linear regression to adjust for large-scale climate processes (see *SI Appendix*), other more local factors that are correlated with river discharge could also be at play. Downwelling-favorable along-shore winds coinciding with anomalously strong river discharge would effect depth-dependent Ekman transports and enhanced coastal trapping of the buoyant outflow, leading to higher coastal sea-level anomalies than would have occurred otherwise (20, 22). Also, we adopted the coastal demarcations of Meade and Emery (25), assuming that these regions are distinct and no communication takes place between them. The observed spatial cross-covariance structure between river-discharge time series (*SI Appendix*, Fig. S2) shows that these regions generally correspond to areas of coherent river-discharge fluctuation. However, instances of significant correlation between rivers from different regions are also seen. Future studies could consider alternative delineations to determine their influence on the relation between river discharge and sea level. Additional factors ignored here—such as background flows, upstream boundary conditions, groundwater discharge, incomplete data records, and the detailed distribution of rivers within any given region—could also play a role. In any case, the general correspondence between theory and data suggests that Eq. 6 encapsulates the most basic and important physical controls (discharge, rotation, and stratification) mediating the relation between annual river-discharge and sea-level anomalies at the coast.

To suggest in more detail the influence of particular rivers on sea level at specific coastal locations, we compare individual river-discharge and sea-level records. Fig. 3 shows a matrix of correlation coefficients between all possible pairs of river-discharge and sea-level time series. The correlation between river discharge and downstream sea level tends to decrease with increasing separation. For most river-discharge records, downstream tide gauges closer to the river are usually more strongly and significantly correlated with the discharge time series than

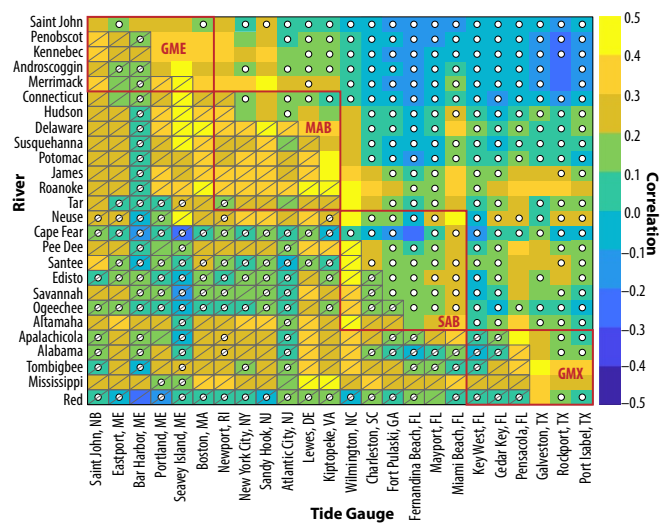


Fig. 3. Color shading shows correlation coefficients between all possible pairs of river discharge and sea level. All time series have been adjusted for large-scale climate modes as with the regional time series and described in *SI Appendix*. White dots indicate values not statistically significant at the $p < 0.05$ level. Grid cells without hatching are downstream sea-level sites, whereas hatched cells indicate upstream tide-gauge sites. Red boxes encapsulate groups of river stations and tide gauges within the same coastal region: GME (Gulf of Maine), MAB (Mid-Atlantic Bight), SAB (South Atlantic Bight), and GMX (Gulf of Mexico).

are tide gauges farther away down the coast. Discharge from the Potomac River is significantly correlated with sea level at the two closest downstream tide-gauge stations at Kiptopeke Beach and Charleston but not significantly correlated with any other downstream sea-level records. However, this general pattern is not universal, and some rivers are exceptions to this rule. The Ogeechee River does not show significant correlation with any downstream tide gauges. It should be noted that this river has relatively small discharge, and its signal in the tide-gauge data could be overcome by the noise due to other local forcing factors.

There are also some interesting, statistically significant correlations between river-discharge records and sea level at upstream tide gauges. Discharge from rivers within Chesapeake Bay (Susquehanna, Potomac, James, Roanoke) are significantly correlated with sea level at sites upstream along the New England coast (Portland, Seavey Island, Boston, Newport). Similarly, there are instances of more remote significant correlations, such as between discharge from the Roanoke River (in Virginia along the Atlantic coast) and sea level at Pensacola, Galveston, Rockport, and Port Isabel (in Florida and Texas along the Gulf coast). It is unlikely that these upstream and remote correlations indicate causation. Rather, they probably reflect underlying correlation between river flows within the same large-scale drainage basins (*SI Appendix*, Fig. S2). For instance, the correlation between outflows from the Roanoke and Apalachicola Rivers is 0.53 (*SI Appendix*, Fig. S2), and the latter river is immediately upstream of the four Florida and Texas tide gauges just mentioned.

Discussion

We investigated the relation between river discharge and sea level along the United States Atlantic and Gulf coasts on inter-annual and longer time scales during 1910–2017. We updated earlier calculations made by Meade and Emery (25) for 1931–1969, demonstrating that their results hold for longer and more recent time periods, and also addressed some outstanding issues prompted by that earlier analysis. We showed that river discharge can be a significant causal driver of year-to-year coastal sea-level

changes and argued that observed correspondences between the two variables are roughly consistent, to order of magnitude, with expectations from theories for alongshore transport downstream of coastal river plumes (18, 24). While its relation to coastal sea level is statistically significant, river discharge explains only a portion of the overall variance in the tide-gauge records. This is unsurprising, given the myriad other processes that influence sea level along this coastline (6–14). Our results thus complement previous sea-level studies, adding to our understanding of the complex and multifaceted nature of sea-level variation along this densely populated coastline.

Although we focused on a particular coastline, our theory is sufficiently general that it should apply to other global coastal ocean locations as well. Applied to the world's 25 largest rivers by annual mean discharge (30), Eq. 5 predicts that river discharge can raise the background mean downstream coastal sea level by ~ 10 cm and can force interannual sea-level variations of $\sim \pm 5$ cm (*SI Appendix, Fig. S3*). The theory should be useful in studies of other areas where river runoff may be an important driver of coastal sea level (31).

Our results have important implications for monitoring coastal-zone changes from space (32). A useful order-of-magnitude length scale characterizing the coastal trapping and offshore decay of these sea-level signals is the baroclinic Rossby deformation radius $L_d = (g'h)^{1/2}/f$, which, by virtue of Eq. 4, can be expressed:

$$L_d = \left(\frac{2g\beta S_0 Q_F}{f^3} \right)^{1/4}. \quad [7]$$

Using the parameter values above and river-discharge data for the United States Atlantic and Gulf coasts, we find that baroclinic Rossby deformation radii are $\lesssim 10$ km. (This argument ignores the influence of shallow bathymetry. If, as in the case of the Mississippi River, coastal-ocean depths are shallower than the layer depth predicted by Eq. 4, which is typically of order 1 to 10 m, offshore length scales of the coastal sea-level anomaly will be broader than L_d (18, 19).) This suggests that coastal sea-level anomalies caused by river discharge may not be captured by conventional satellite-altimetry data products, which have poor resolution within 10 to 20 km of the coast, due to uncertainties in correction algorithms and contamination of the satellite footprint (33). This could partly explain the poor correspondence between interannual sea level from tide gauges and satellite altimetry observed in some coastal regions (34). These results underscore the crucial importance of ongoing coastal-altimetry data reprocessing efforts (35) and future high-resolution satellite-altimetry missions (36) for monitoring the coastal ocean to provide empirical linkages between the terrestrial and marine realms.

Our findings show that the influence of river discharge on coastal sea level can be regional and nonlocal, extending downstream along the coast for hundreds of kilometers (23) (Fig. 3). This has important implications for the interpretation of tide-gauge records in the context of ocean-circulation and climate studies. Previous authors interpret low-frequency sea-level variations from tide-gauge records in terms of fluctuations in large-scale ocean circulation (28, 37–40). Our findings suggest that such interpretations should be made carefully, such that sea-level time series from locations downstream of major rivers should be used with caution in this context. This point raises interesting questions, which should be taken up in later studies. For example, what is the ultimate fate of buoyant coastal outflows, and how far down the coast can their influence be observed? Relatedly, note that we focused on coastal sea level along open-ocean boundaries and in the far-field downstream of rivers. We did not consider sea level at more near- or midfield

sites within bays, estuaries, or river mouths. Since many populous United States cities are situated directly on major rivers (e.g., Philadelphia, Baltimore, Washington, DC), future efforts are needed to understand the relation between river discharge and sea level in this latter case, since the basic dynamical balances (e.g., hydraulic control, gradient-wind balance) are expected to be distinct from those considered here (24).

These results have clear implications for modeling and projection. To accurately simulate future discharge-driven sea-level changes and their coastal impacts, climate models would need to transfer accurately projected precipitation changes over a drainage basin to the coastal ocean and use sufficiently high spatial resolution over the ocean to resolve the attendant dynamical response. Most Coupled Model Intercomparison Project Phase 5 (CMIP5) models use horizontal ocean resolutions on the order of ~ 10 to 100 km (41), which is too coarse to resolve discharge-driven coastal sea-level changes. Thus, regional-scale flood risk in a given year may not be accurately represented in these projections. It will take another decade of model development before state-of-the-art global climate models have sufficiently high horizontal resolution to resolve the first baroclinic Rossby radius over most of the shallow (< 500 m) global ocean (42). This suggests that alternative means of incorporating the discharge effects into model projections (e.g., improved parameterization schemes for impacts at the coast) warrant further attention. Such efforts will be crucial in light of projected future intensification of the hydrological cycle. For example, discharge of the Mississippi River is projected to increase by 11% to 60% over the coming century (43). Eq. 5 suggests that this increase in discharge would cause a background mean coastal sea-level rise downstream of the Mississippi River of 0.5 to 2.6 cm. Such considerations should be factored into centennial coastal sea-level projections at local to regional spatial scales.

Finally, these results have potential implications for interpretation of proxy relative sea-level reconstructions. Derived from ~ 1 cm-thick salt-marsh sediment samples, Common Era proxy sea-level reconstructions are inherently time averaged. Depending on background rates of glacial isostatic adjustment, the temporal averaging implicit in a sediment sample can vary, for example, from ~ 30 y in Florida (44) to ~ 7 y in southern New Jersey (45). To evaluate whether proxy-relative sea-level reconstructions are influenced by river discharge, we extended our analysis by averaging the data over different multiyear time periods (*SI Appendix, Fig. S4*). The correlation between river discharge and sea level generally decreases and becomes statistically insignificant for time scales $\gtrsim 5$ to 10 y on the Atlantic coast and $\gtrsim 10$ y along the Gulf Coast (*SI Appendix, Fig. S4*). While this finding could imply that river discharge is not an important contributor to the decadal and longer sea-level variations reflected in proxy reconstructions, it could also reflect the relative shortness of the data records given the time scales relevant to paleo problems. Also possibly important in this context are the broader distances over which advection by the alongshore flow can act on increasingly long time scales. Model studies would therefore be informative in these regards. If river discharge exerts an appreciable control on sea level at decadal and longer time scales, its influence on proxy reconstructions would be regional in scale (i.e., not restricted to sites in or very close to estuaries but also present at far-field sites along the open coast), as evidenced by the spatial scales of correlation seen in the instrumental record.

ACKNOWLEDGMENTS. C.G.P. and R.M.P. acknowledge support from NASA Contract NNH16CT01C (which also supported C.M.L.), NASA Jet Propulsion Laboratory Subcontract 1569246, and National Science Foundation Award 1558966. C.G.P. also acknowledges support from The Investment in Science Fund at Woods Hole Oceanographic Institution. A.C.K. and S.E.E. acknowledge NSF Awards OCE-1458921 and OCE-1458903, respectively. This is a contribution to the International Geoscience Programme Project 639 “sea-level change from minutes to millennia” and PALSEA 2 (paleo constraints on sea level rise 2).

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